

NUCLEATION IN SYNOPTICALLY FORCED CIRRUS

Ruei-Fong Lin

University of Maryland Baltimore County

David O'C. Starr

NASA Goddard Space Flight Center

Jens Reichardt

University of Maryland Baltimore County

Paul J. DeMott

Colorado State University

Cold Cirrus

T < -40°C

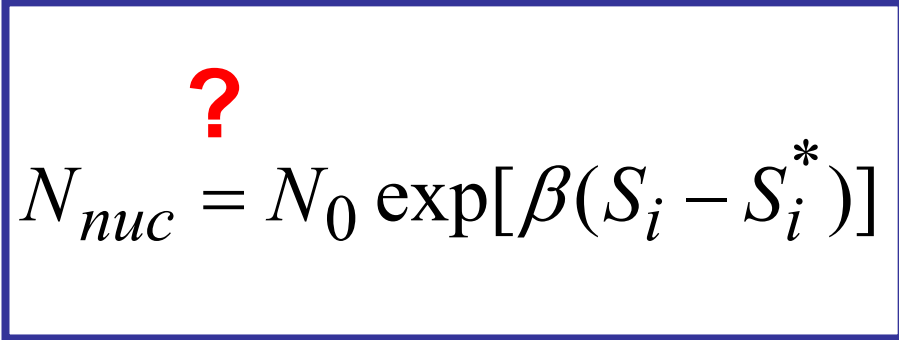
Following Meyers et al. (1992)

Homogeneous Nucleation

- Homogeneous freezing of concentrated aqueous solution droplets at water subsaturated conditions.

Heterogeneous Nucleation

- Activation of IN in concentrated aqueous solution droplets at water subsaturated conditions.
- Deposition nucleation


$$N_{nuc} = N_0 \exp[\beta(S_i - S_i^*)]$$

$$N_0 = 1 \text{ liter}^{-1}$$

S_i : supersaturation ratio wrt. ice

$$\beta = \beta(T, w, \text{IN species})$$
$$S_i^* = S_i^*(T, w, \text{IN species})$$

$$N_{nuc} = N_0 \exp[\beta(S_i - S_i^*)]$$

Meyers et al (1992): Deposition-Condensation Freezing

$$-7 > T > -20^\circ\text{C}, 0.25 > S_i > 0.02$$

$$\beta \sim 12.96, S_i^* \sim 0.0493 \text{ (hereafter, M92)}$$

Immersion freezing

Parcel Model Studies:

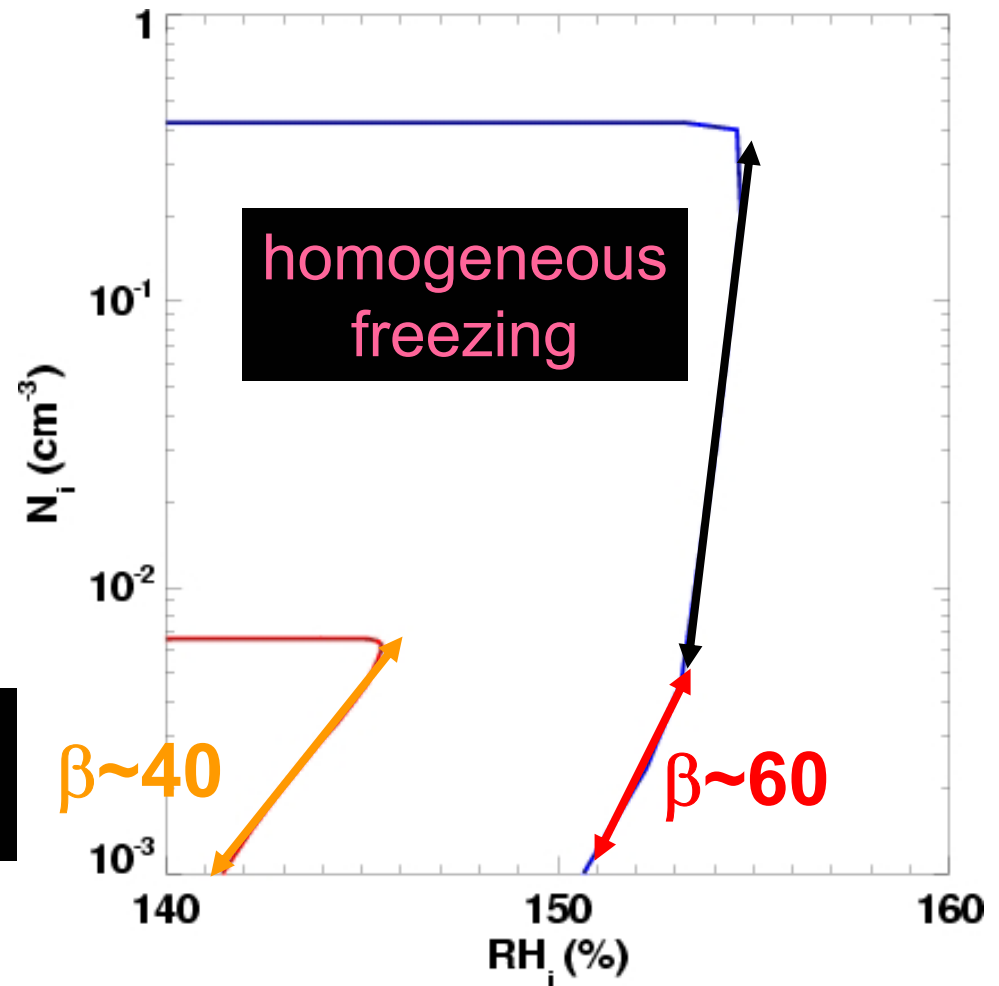
$$w = 0.04 \text{ m s}^{-1}$$

$$T = -60^\circ\text{C}$$

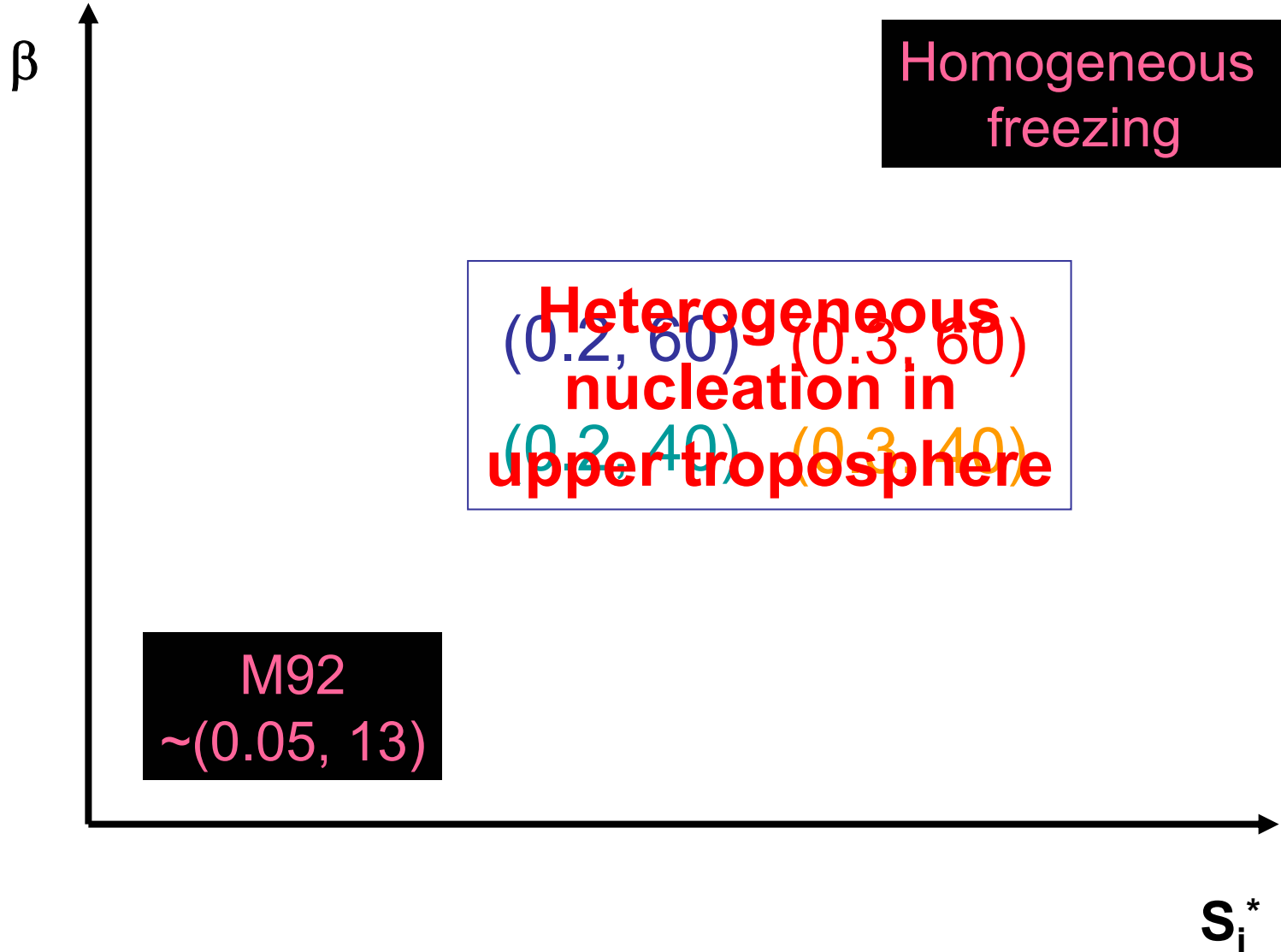
$$T = -40^\circ\text{C}$$

(DeMott)

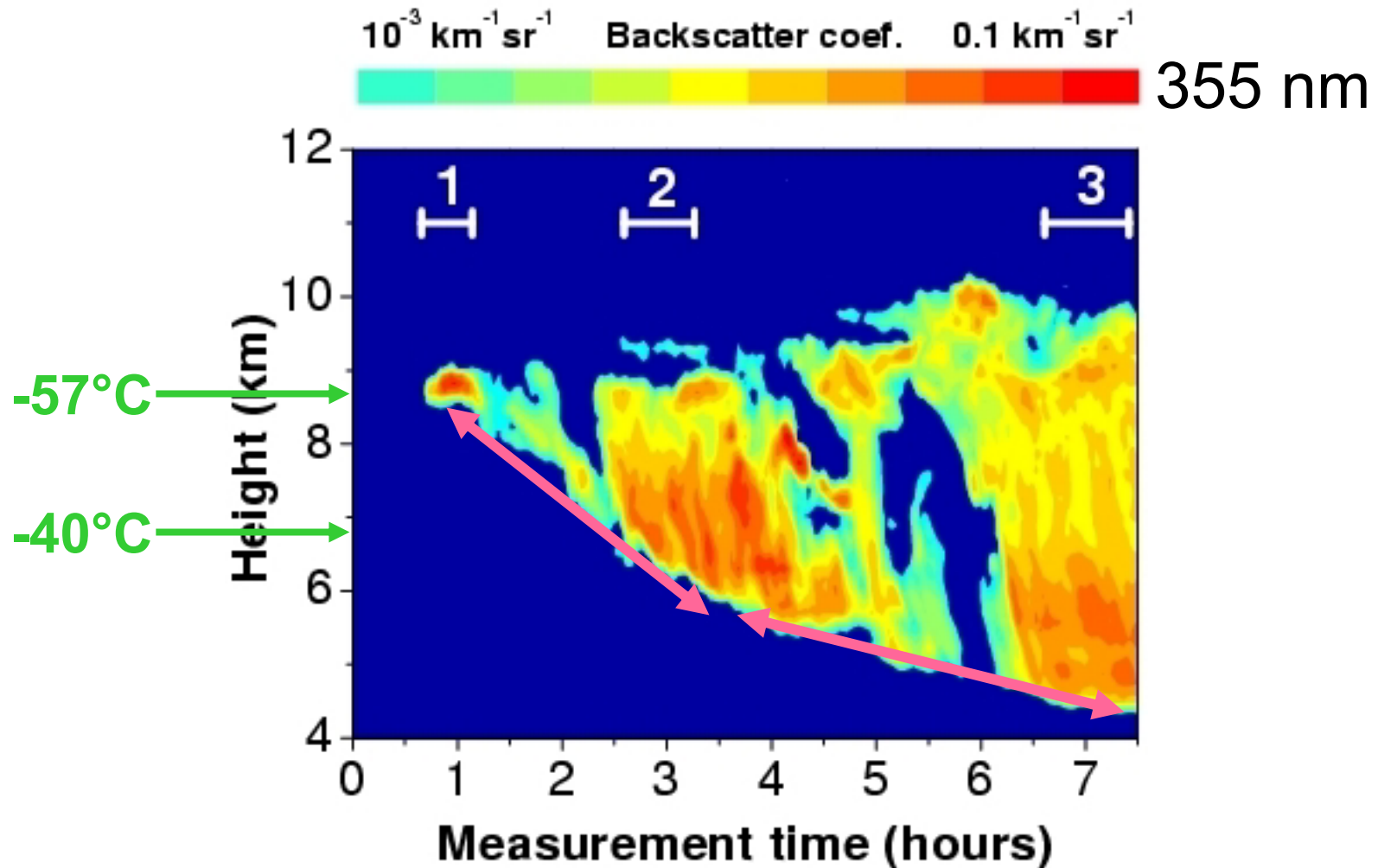
immersion
freezing



$$(S_i^*, \beta)$$



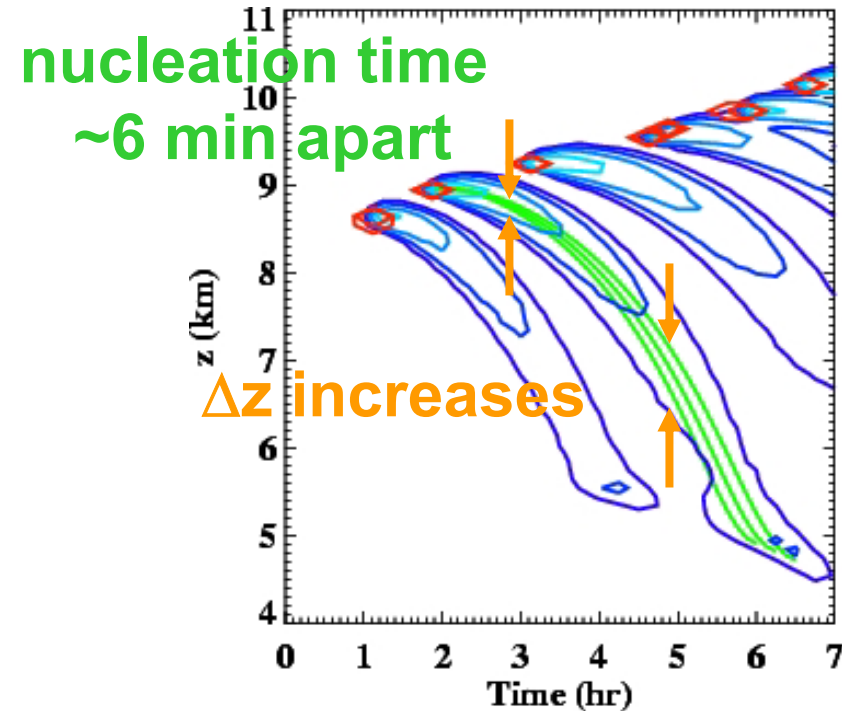
Cirrus cloud measured over northern Sweden.



Reichardt et al. (2002), GRL, **29**, 10.1029/2002GL014836.

1D Cirrus Simulation with Explicit Microphysics

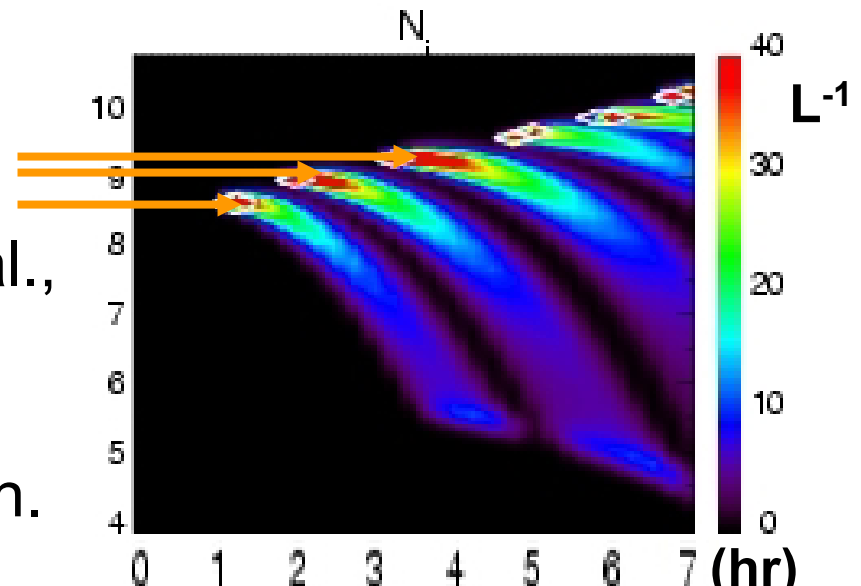
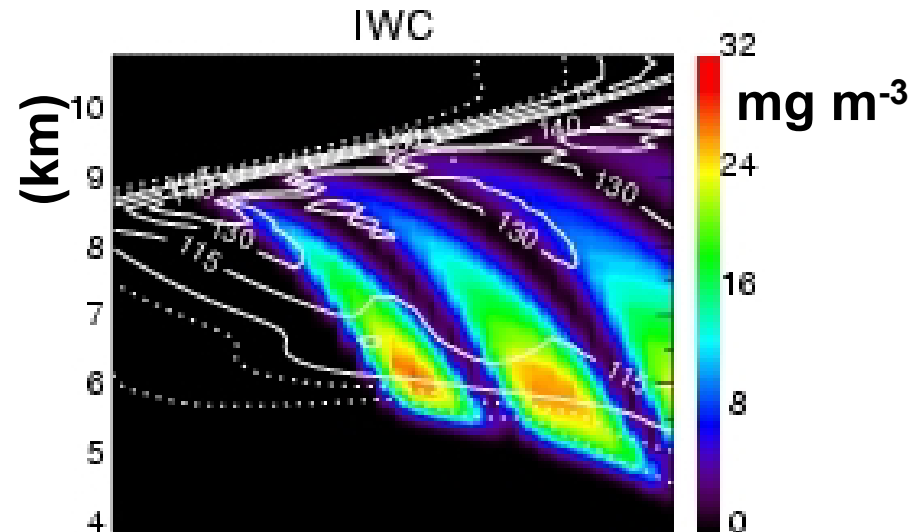
Ice particle trajectories



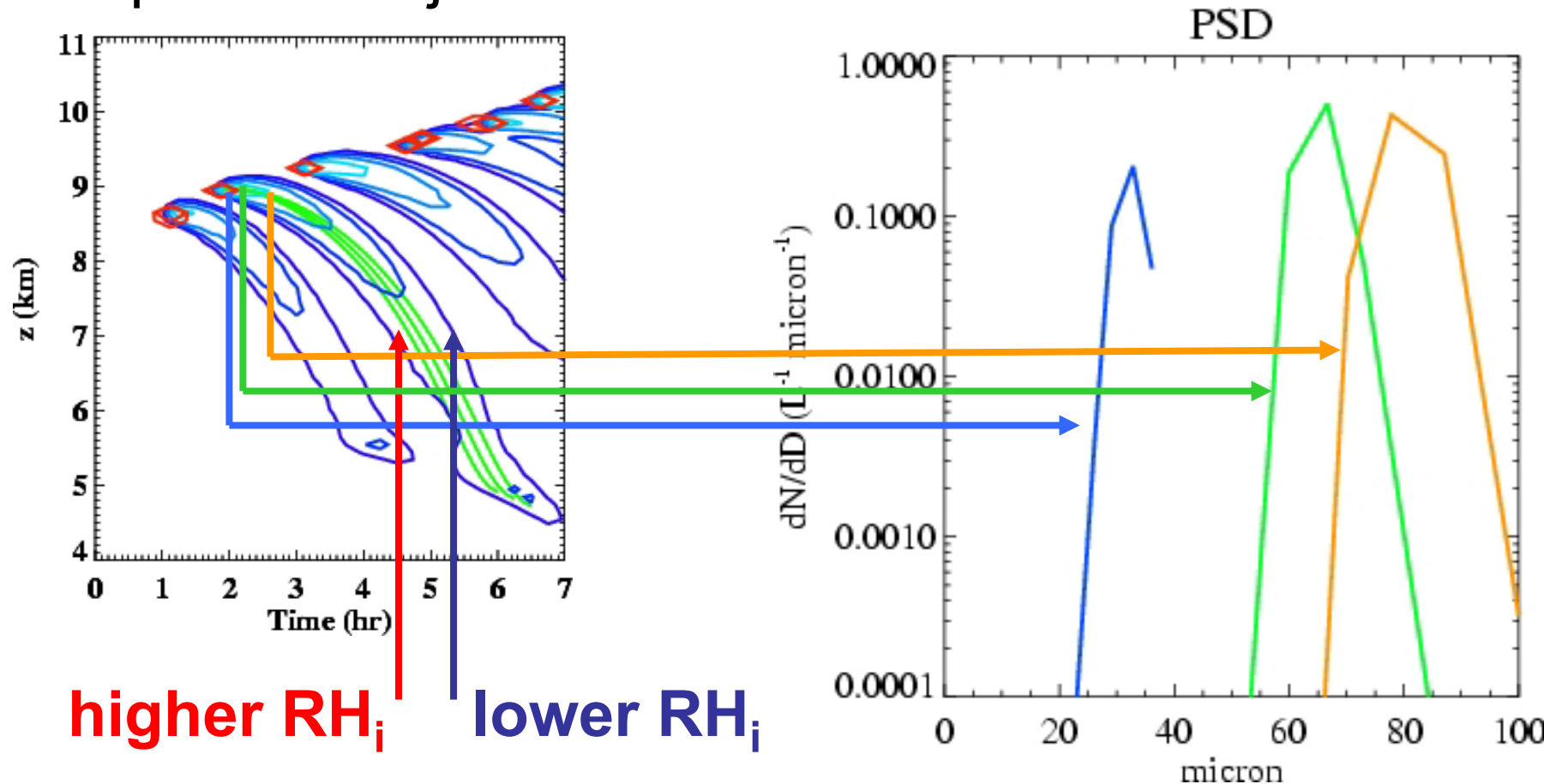
Nucleation pulse (Sassen and Dodd, 1989; Khvorostyanov et al., 2001) at the very top of each cloud episode.

Nucleation lasts for about 15 min.

Homogeneous freezing



Ice particle trajectories



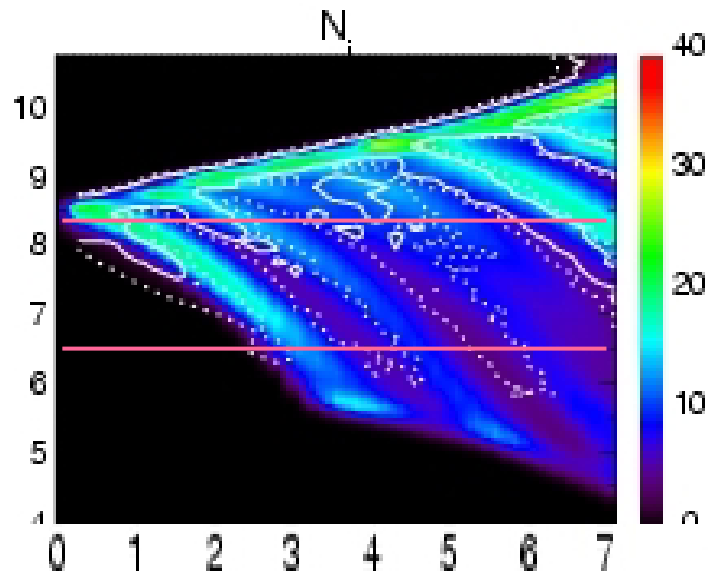
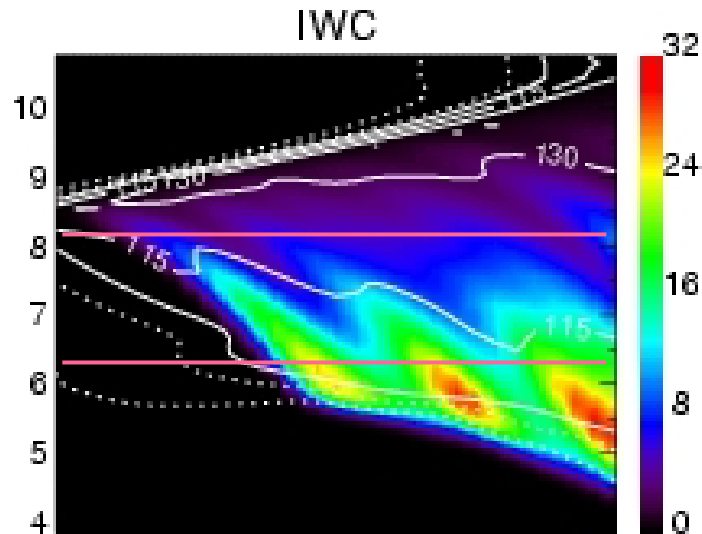
higher RH_i **lower RH_i**

LONG NUCLEATION DURATION

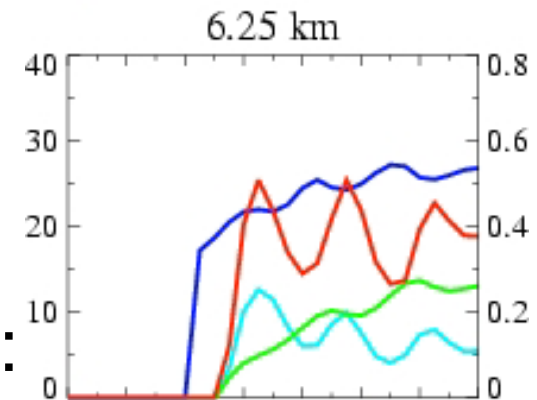
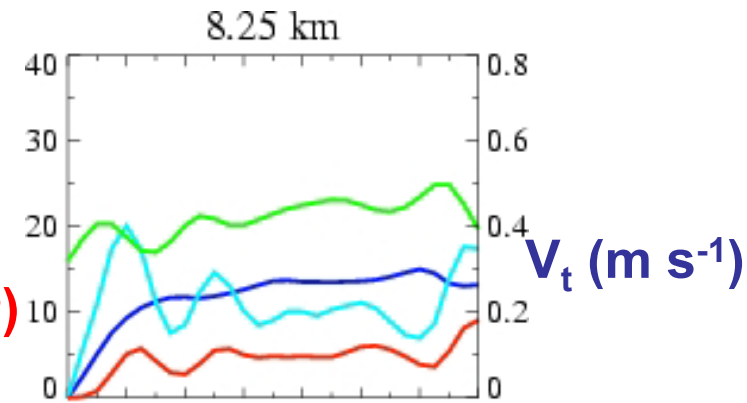
N_i generated at the cloud top
does not provide sufficient information of
 N_i in the mass bearing regime

1D Cirrus Simulation with Explicit Microphysics

Heterogeneous Nucleation M92



RH_i -100%
IWC (mg m^{-3})
 N_i (L^{-1})

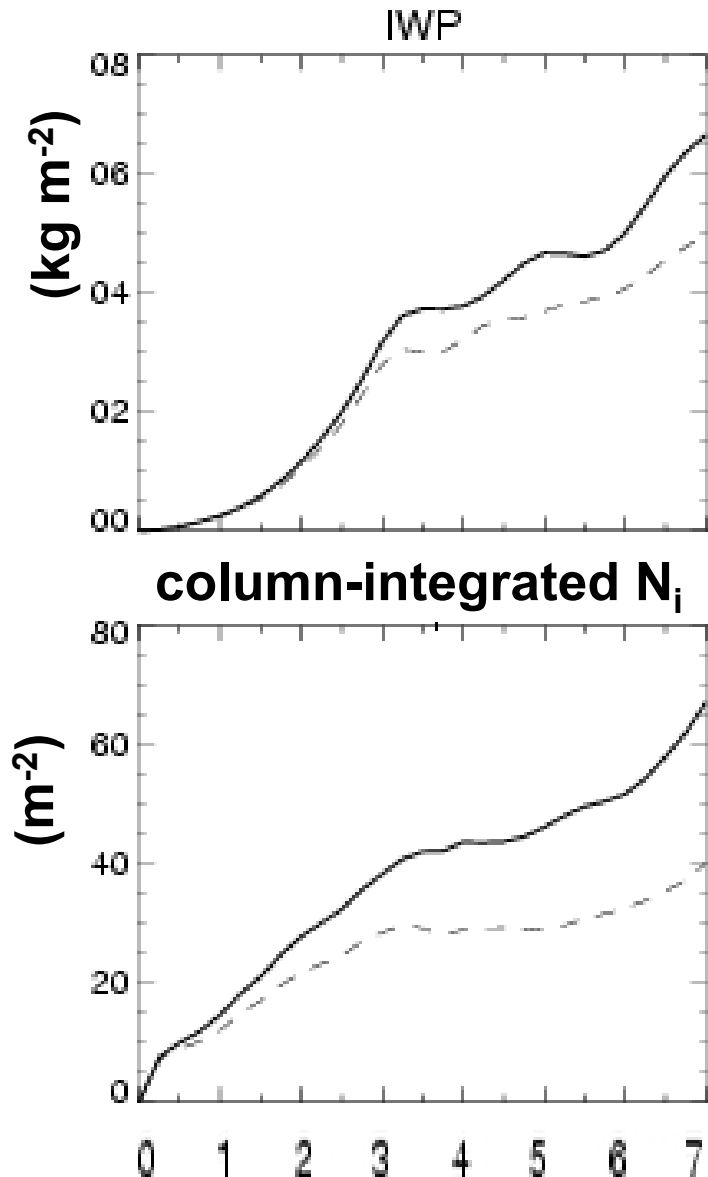


Nucleation:
Cloud top: continuous, small variation with time
Mid-cloud: intermittent, active when RH is on the rise and IN are available

Average cloud properties are sensitive to mid-cloud nucleation

solid: control run

**dashed: mid-cloud
nucleation
switched off**



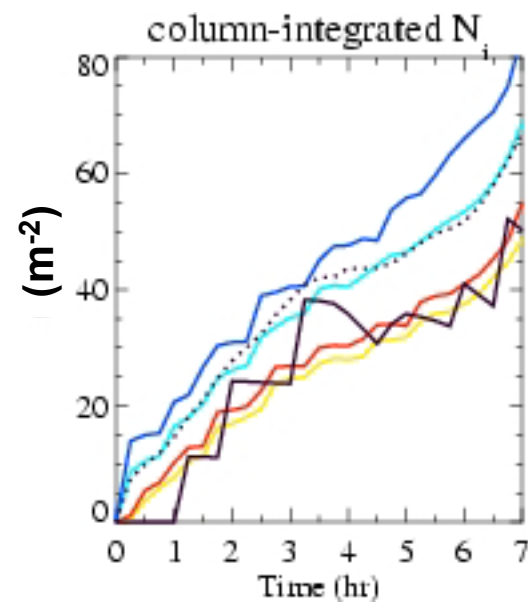
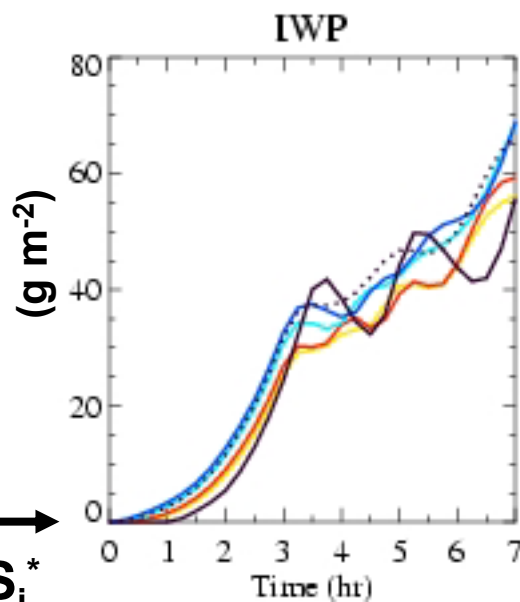
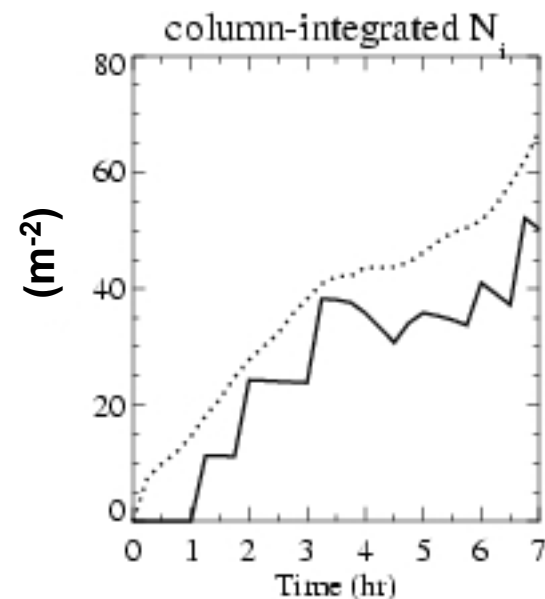
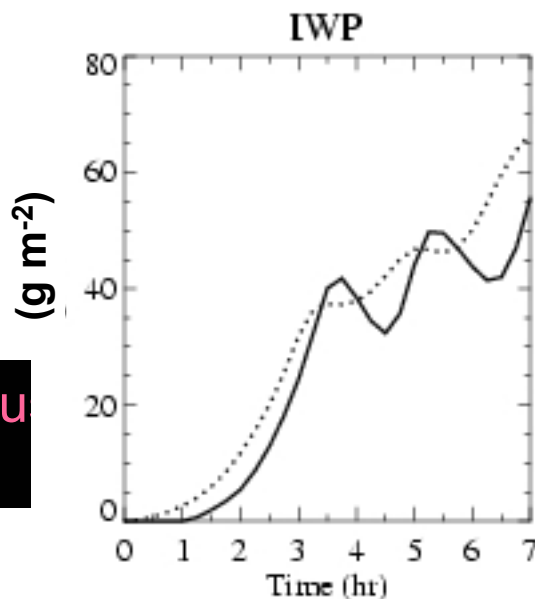
(S_i^*, β)

Homogeneous
freezing

(0.2, 60) (0.3, 60)

(0.2, 40) (0.3, 40)

M92
~(0.05, 13)

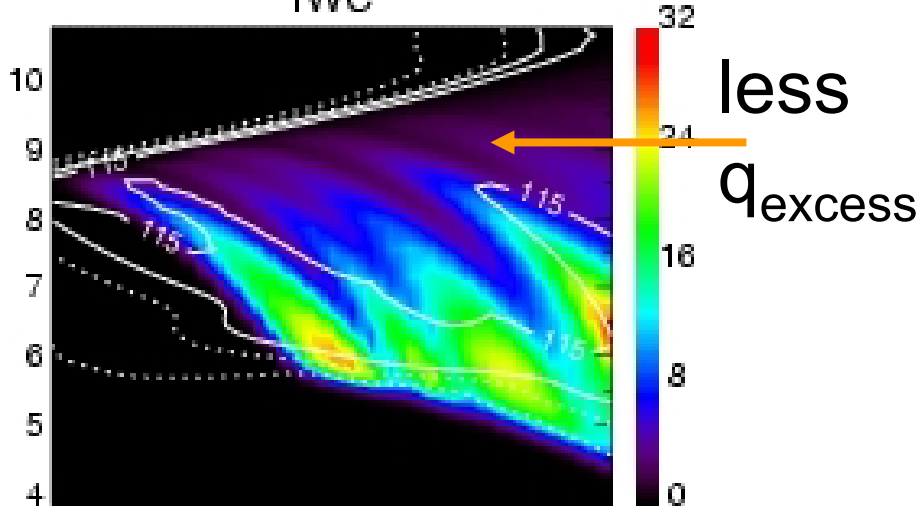


S_i^*

(\mathbf{S}_i^*, β)

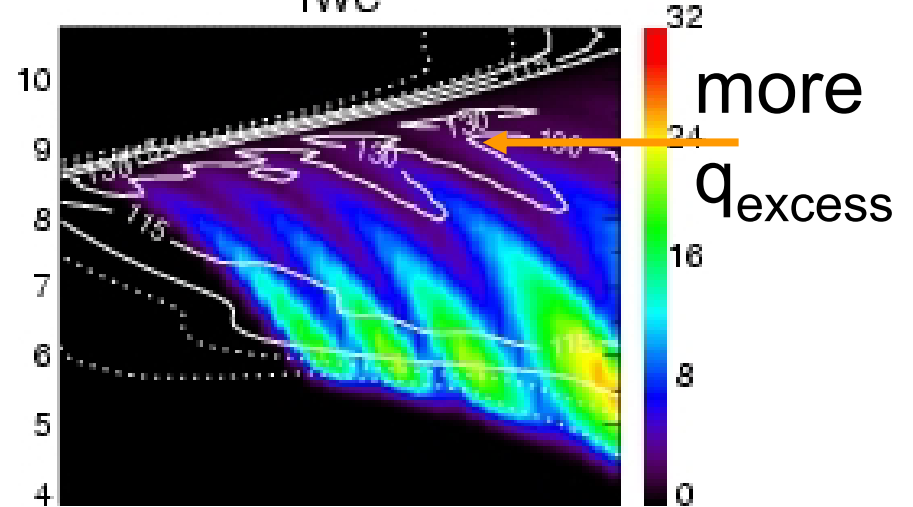
$(0.2, 60)$

IWC

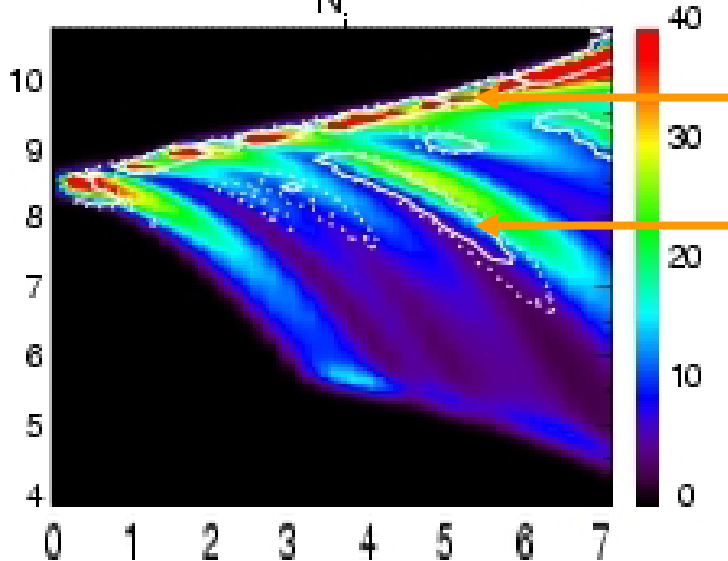


$(0.3, 60)$

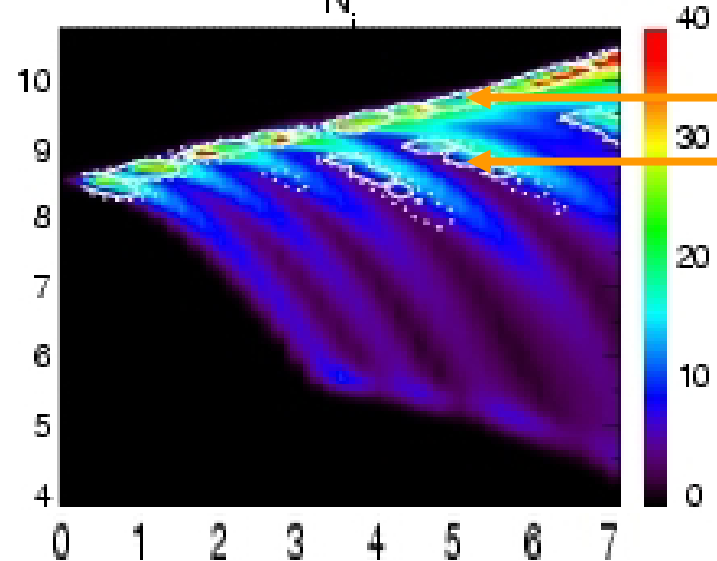
IWC



N_i



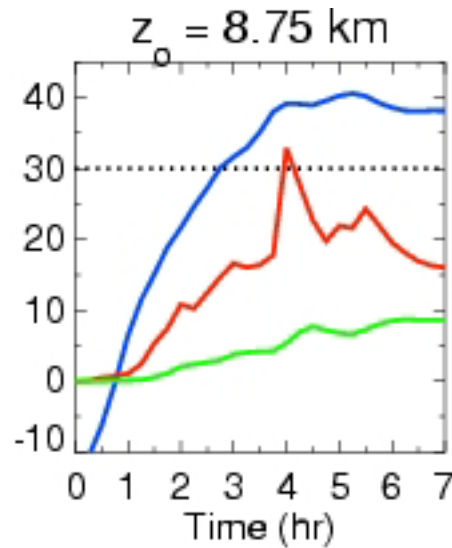
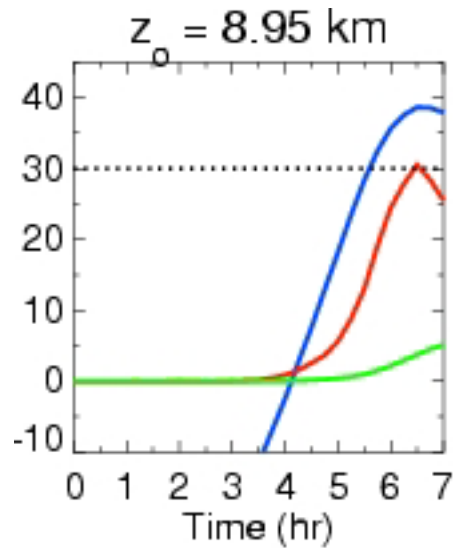
N_i



Summary

- **Synoptically forced cirrus: long duration of nucleation period.** N_i at nucleation zones does not provide sufficient information of N_i in the cloud mass bearing regions.
- Preliminary results show that **the average cloud properties do not respond to S_i^* and β monotonically.** Sensitivity tests must cover a reasonable “2-dimensional” range of these two parameters.
- For the limited range of (S_i^*, β) studied here, the average cloud properties are more sensitive to S_i^* . Smaller S_i^* entails less excess water vapor, more cloud ice, larger nucleation zone, and more ice crystals.

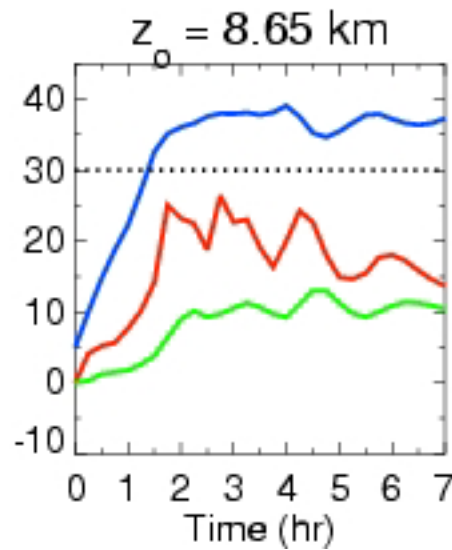
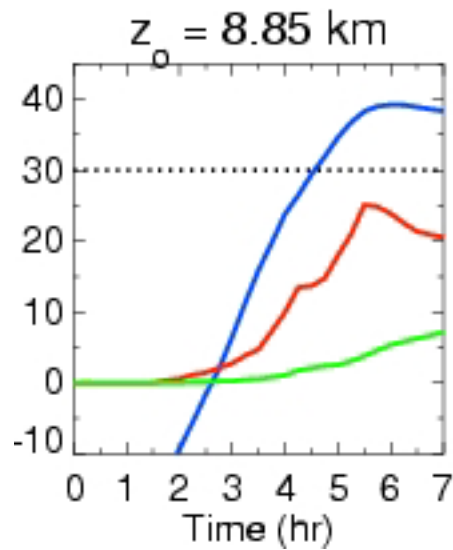
Following ascending parcels



$RH_i - 100\%$

N_i (L^{-1})

IWC ($mg\ m^{-3}$)



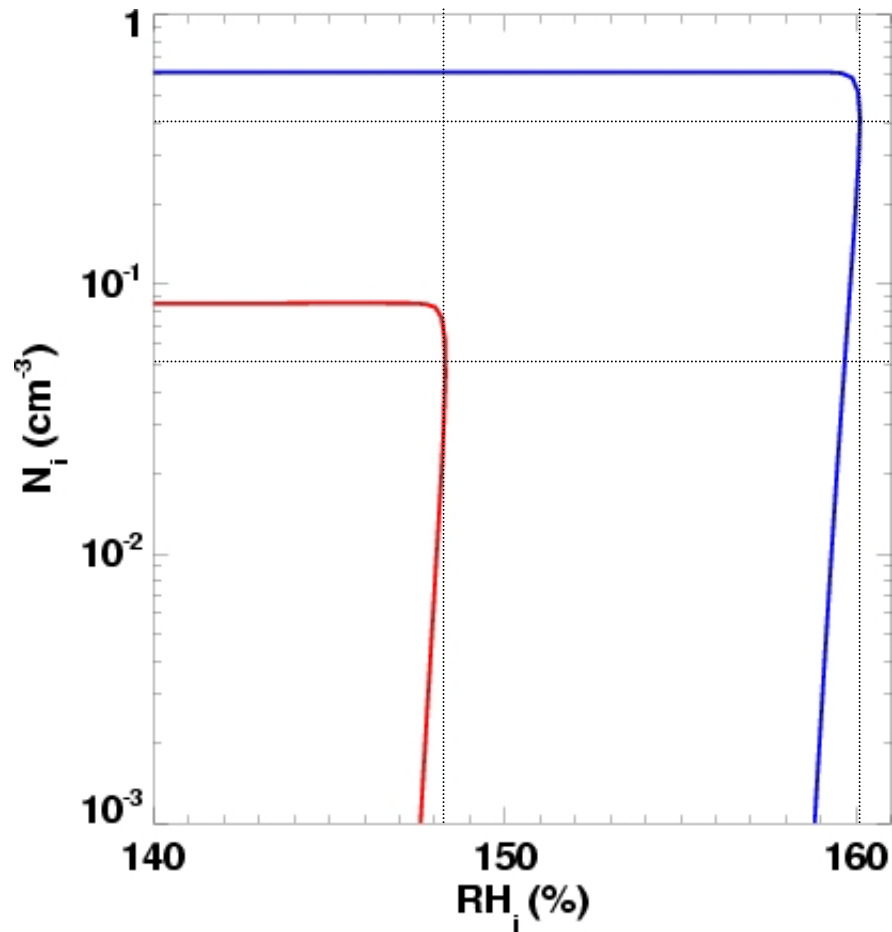
$$N_{nuc} = N_0 \exp[\beta(S_i - S_i^*)]$$

Homogeneous freezing (from Parcel Model Studies)

$w = 0.04 \text{ m s}^{-1}$

$T = -60^\circ\text{C}$

$T = -40^\circ\text{C}$



Caveat: May underestimate N_{nuc}